Boundary conditions control in ORCA2

Eugene Kazantsev

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Journée thématique
"Que peuvent attendre les modélisateurs de l'assimilation de données ?"
Paris, le 12 février 2013

ORCA-2 configuration of NEMO

- $182 \times 149 \times 31$ nodes in curvilinear (x, y) coordinates;
- z levels with partial steps at the bottom;
- leap-frog scheme with Asselin filter;
- implicit surface pressure gradient with External Gravity Waves filter;
- implicit vertical diffusion with TKE Turbulent Closure Scheme;
- Solar Radiation + Geothermal Heating + BBL + Surface evaporation/precipitation;
- surface wind stress.

Space discretization on the C-grid

$$
\begin{aligned}
\frac{\partial u}{\partial t} &= \Big(S_x S_y v\Big) S_y(\omega + f) - D_x \frac{S_x u^2 + S_y v^2}{2} - S_z\Big(S_x w D_z u\Big) + D_x A_u^h \xi + D_y A_u^h \omega + \\
&+ g \int_0^z D_x S_z \rho(x,y,\zeta) d\zeta + D_{zz}(A_u^z u) + g D_x(\eta + T_c \phi) \\
\frac{\partial T}{\partial t} &= -D_x(u S_x T) - D_y(v S_y T) - D_z(w S_z T) + A_T^h \Big(D_x D_x T + D_y D_y T\Big) + \\
&+ D_{zz}(A_T^z T) + \text{Solar Radiation} + \text{Geothermal Heating} + \text{BBL} \\
\xi &= D_x u + D_y v, \quad \omega = D_y u - D_x v, \quad w = \int_H^z \xi(x,y,\zeta) d\zeta; w(x,y,H) = 0
\end{aligned}
$$

Interpolations and Derivatives

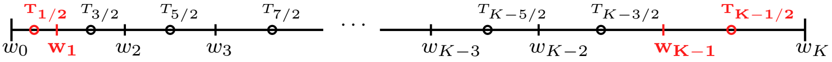
$$
\begin{aligned}
(Sw)_{k+1/2} &= \frac{w_{k+1} + w_k}{2} \; k = 1, \ldots, K-1 \\
(DT)_k &= \frac{T_{k+1/2} - T_{k-1/2}}{h} \; k = 1, \ldots, K-1
\end{aligned}
$$

Space discretization on the C-grid

$$
\begin{aligned}
\frac{\partial u}{\partial t} &= \Big(S_x S_y v\Big) S_y(\omega + f) - D_x \frac{S_x u^2 + S_y v^2}{2} - S_z \Big(S_x w D_z u\Big) + D_x A_u^h \xi + D_y A_u^h \omega + \\
&+ g \int_0^z D_x S_z \rho(x,y,\zeta) d\zeta + D_{zz}(A_u^z u) + g D_x(\eta + T_c \phi) \\
\frac{\partial T}{\partial t} &= -D_x(u S_x T) - D_y(v S_y T) - D_z(w S_z T) + A_T^h \Big(D_x D_x T + D_y D_y T\Big) + \\
&+ D_{zz}(A_T^z T) + \text{Solar Radiation} + \text{Geothermal Heating} + \text{BBL} \\
\xi &= D_x u + D_y v, \quad \omega = D_y u - D_x v, \quad w = \int_H^z \xi(x,y,\zeta) d\zeta; w(x,y,H) = 0
\end{aligned}
$$

Interpolations and Derivatives Modified Near the boundary

$$
\begin{aligned}
(Sw)_{k+1/2} &= \frac{w_{k+1} + w_k}{2}, \ k = 1, \dots, K-2, \ (Sw)_{1/2} = \alpha_0^S + \alpha_1^S w_0 + \alpha_2^S w_1 \\
(DT)_k &= \frac{T_{k+1/2} - T_{k-1/2}}{h}, \ i = 2, \dots, K-2, \ (DT)_1 = \alpha_0^D + \frac{\alpha_1^D T_{1/2} + \alpha_2^D T_{3/2}}{h}
\end{aligned}
$$

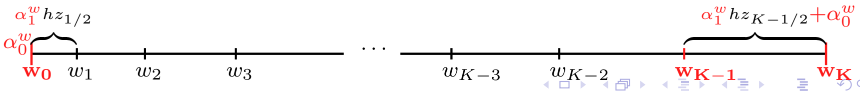


Space discretization on the C-grid

$$
\begin{aligned}
\frac{\partial u}{\partial t} &= \Big(S_x S_y v\Big) S_y(\omega + f) - D_x \frac{S_x u^2 + S_y v^2}{2} - S_z \Big(S_x w D_z u\Big) + D_x A_u^h \xi + D_y A_u^h \omega + \\
&+ \quad g \int_0^z D_x S_z \rho(x,y,\zeta) d\zeta + D_{zz}(A_u^z u) + g D_x(\eta + T_c \phi) \\
\frac{\partial T}{\partial t} &= -D_x(u S_x T) - D_y(v S_y T) - D_z(w S_z T) + A_T^h \Big(D_x D_x T + D_y D_y T\Big) + \\
&+ \quad D_{zz}(A_T^z T) + \text{Solar Radiation} + \text{Geothermal Heating} + \text{BBL} \\
\xi &= D_x u + D_y v, \quad \omega = D_y u - D_x v, \quad w = \int_H^z \xi(x,y,\zeta) d\zeta; w(x,y,H) = \alpha_0(x,y)
\end{aligned}
$$

Vertical velocity

$$
\begin{aligned}
w_{i,j,K-1} &= \alpha_0^{w^b} - \alpha_1^{w^b} hz_{i,j,K-1/2}\xi_{i,j,K-1/2} \\
w_{i,j,k-1} &= w_{i,j,k} - hz_{i,j,k-1/2}\xi_{i,j,k-1/2} \quad \forall k : 2 \le k \le K-1 \\
w_{i,j,0} &= w_{i,j,1} + \alpha_0^{w^s} - \alpha_1^{w^s} hz_{i,j,1/2}\xi_{i,j,1/2}
\end{aligned}
$$



Vertical diffusion

$$\frac{\partial}{\partial z} A_u^z \frac{\partial u}{\partial z} \quad \text{is replaced by}$$

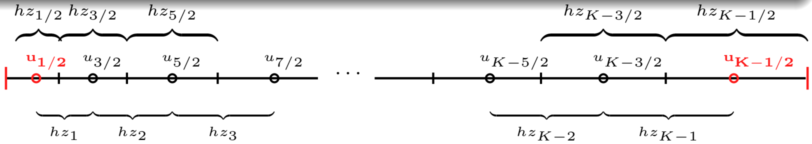
$$\left(D_{zz}u\right)_{i,j,1/2} = \frac{(A_u^z)_1}{hz_1 hz_{1/2}}(\alpha_2^{DzzU^s} u_{3/2} - \alpha_1^{DzzU^s} u_{1/2})$$

$$\left(D_{zz}u\right)_{i,j,k-1/2} = \frac{1}{hz_{k-1/2}}\left(\frac{(A_u^z)_k}{hz_k}(u_{k+1/2} - u_{k-1/2}) - \frac{(A_u^z)_{k-1}}{hz_{k-1}}(u_{k-1/2} - u_{k-3/2})\right) \quad \forall k : 2$$

$$\left(D_{zz}u\right)_{i,j,K-1/2} = \frac{1}{hz_{K-1/2}}\left[\alpha_2^{DzzU^b} \frac{(A_u^z)_{K-1}}{hz_{K-1}} u_{K-1/2} - \alpha_1^{DzzU^b}\left(\frac{(A_u^z)_K}{hz_K} + \frac{(A_u^z)_{K-1}}{hz_{K-1}}\right) u_{K-3}\right.$$

$$\left.\frac{\partial u}{\partial z}\right|_{w0} = \alpha_0^{DzzU^s} + \frac{\tau_x}{hz_1 \rho_0}, \quad \left.\frac{\partial v}{\partial z}\right|_{w0} = \alpha_0^{DzzU^s} + \frac{\tau_y}{hz_1 \rho_0}, \quad \left.\frac{\partial T}{\partial z}\right|_{w0} = \left.\frac{\partial S}{\partial z}\right|_{w0} = \alpha_0^{DzzT^s}$$

$$u|_{bottom} = v|_{bottom} = \alpha_0^{DzzU^b} \quad T|_{bottom} = S|_{bottom} = \alpha_0^{DzzT^b} \tag{1}$$



The models solution depend on initial and boundary conditions :

$$\frac{\partial T}{\partial t} = -D_x(uS_xT) - D_y(vS_yT) - D_z^{(\alpha)}(wS_z^{(\alpha)}T) + A_T^h\left(D_{xx}T + D_{yy}T\right) + Dzz^{(\alpha)}(A_T^z)T$$

- The model $x(t) = \mathcal{M}_{0,t}(x_0, \alpha)$

We calculate the derivatives and their adjoints with respect to

$$x_0, \quad \alpha$$

by TAPENADE 3.6 (Tropics team, INRIA) that allows us

- to avoid a HUGE development/coding (a double of the classical one, at least)
- to obtain immediately the derivative with respect to any parameter we want.

TAPENADE 3.6 (Tropics team, INRIA) with the Memory Usage Optimization:

search for push/pop

```
CALL PUSHREAL8ARRAY(sold, nx*ny*nz)
CALL PUSHREAL8ARRAY(told, nx*ny*nz)
CALL PUSHREAL8ARRAY(vold, nx*ny*nz)
CALL PUSHREAL8ARRAY(uold, nx*ny*nz)
CALL PUSHREAL8ARRAY(ssh, nx*ny)
CALL PUSHREAL8ARRAY(s, nx*ny*nz)
CALL PUSHREAL8ARRAY(t, nx*ny*nz)
CALL PUSHREAL8ARRAY(v, nx*ny*nz)
CALL PUSHREAL8ARRAY(u, nx*ny*nz)
```

replace by

```
call push_uvts(u,v,t,s,ssh)
```

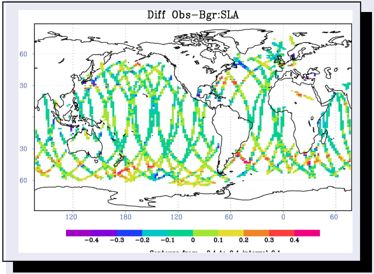
Procedure push/pop_uvts(u,v,t,s,ssh):

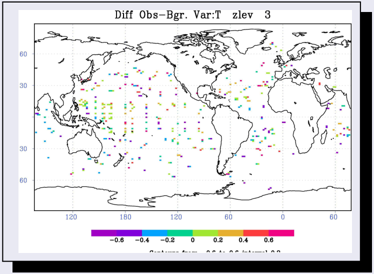
- does not push $n-1$ step and pops appropriate values (divides the required memory by 2)
- does not push u, v, t, s in lower level routines
- does not push values on continents (divides by 2)
- pushes values in Real*4 format (divides by 2)
- eventually pushes only odd timesteps and interpolate when poping (divides by 2)

Total reduction of required memory is up to 25 times.
10 hours window $\Longrightarrow$ 10 days window.

ECMWF data issued from Jason-1 and Envisat altimetric missions and ENACT/ENSEMBLES data banque.

January, 1, 2006.





Difference between observations and background during the 1st of January.

The model: $x_N = \mathcal{M}_{0,N}(x_0, \alpha)$ with $x = (u, v, T, S, ssh)^T$

Cost function J

$$
\begin{aligned}
J &= \|x_0 - x_{bgr}\|^2_{B^{-1}} + \|\alpha - \alpha_{bgr}\|^2_{B^{-1}} + \\
&+ \sum_{n=0}^{N} t_n \|\mathcal{H}\mathcal{M}_{0,n}(x_0, \alpha) - y_n\|^2_{R^{-1}}
\end{aligned}
$$

Matrices: $B^{-1} = diag(10^{-4})$,
$R^{-1} = diag(1/\sigma_u, 1/\sigma_v, 1/\sigma_T, 1/\sigma_S, 1/\sigma_{ssh})$ where $\sigma_u^2 = \frac{1}{N_{obs}} \sum (u_{obs} - u_{bgr})^2$
Minimization is performed by M1QN3 (JC Gilbert, C.Lemarechal)

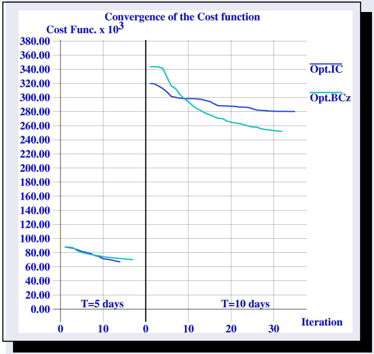
Data Assimilation – Forecast

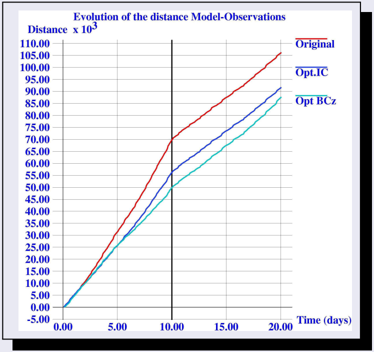
Assimilation window — 10 days (Jan. 1-10, 2006),
Test time — 20 or 30 days (Jan. 1-31, 2006).

The model: $x(t) = \mathcal{M}_{0,t}(x_0, \alpha)$ with $x = (u, v, T, S, ssh)^T$

$$\text{Distance: } \xi(t) = \sum_{n=0}^{t} \| \mathcal{H}\mathcal{M}_{0,n}(x_0, \alpha) - y_n \|_{R^{-1}}$$

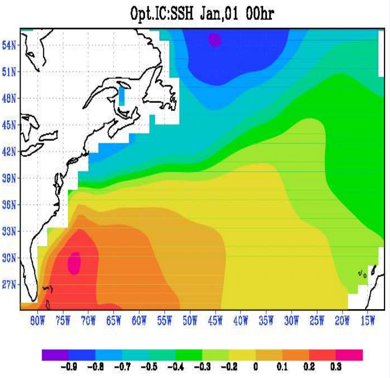
Convergence of J and evolution of ξ



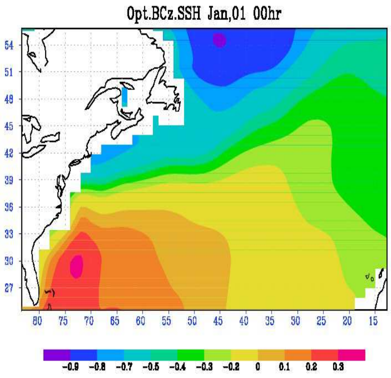


20 Cost function calls with $T = 5$ days and 40 calls with $T = 10$ days.

SSH, North Atlantic, January,1-30 2006.

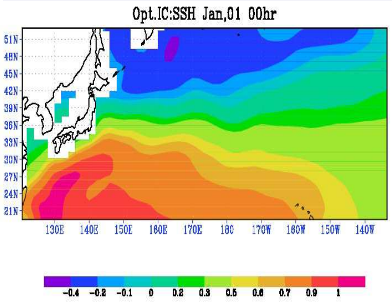


Optimal IC

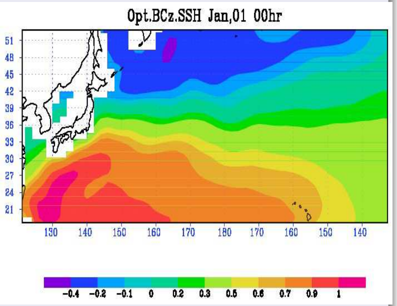


Optimal BCz

SSH, North Pacific, January,1-30 2006.



Optimal IC



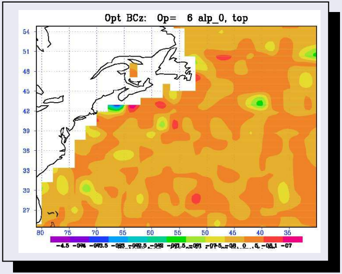
Optimal BCz

Modified formula

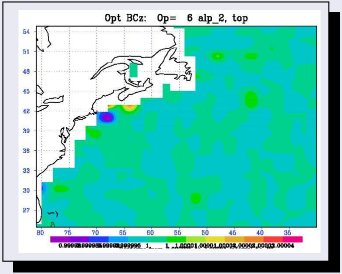
$$
\begin{aligned}
w_{i,j,K-1} &= \alpha_0^{w^b} - \alpha_2^{w^b} hz_{i,j,K-1/2}\xi_{i,j,K-1/2} \\
w_{i,j,k-1} &= w_{i,j,k} - hz_{i,j,k-1/2}\xi_{i,j,k-1/2} \quad \forall k : 1 \le k \le K-2 \\
w_{i,j,0} &= w_{i,j,1} + \alpha_0^{w^s} - \alpha_2^{w^s} hz_{i,j,1/2}\xi_{i,j,1/2}
\end{aligned}
$$

BC for the vertical velocity

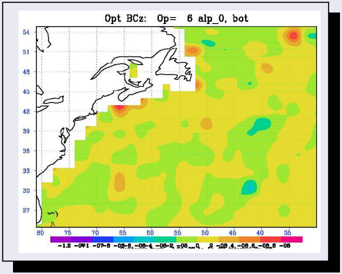
α for the vertical velocity w. North Atlantic.



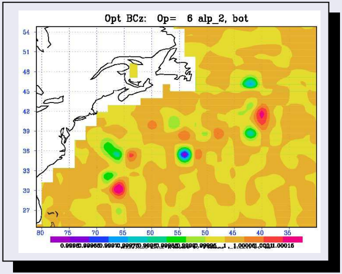
α_0 on the surface



α_2 on the surface

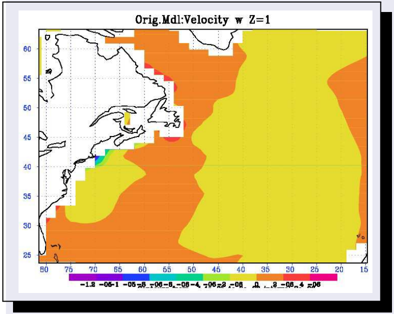


α_0 on the bottom

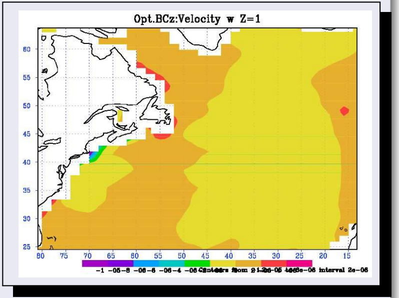


α_2 on the bottom

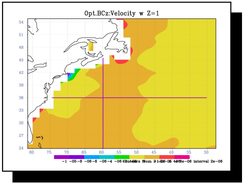
North Atlantic, January, 30, 2006, surface



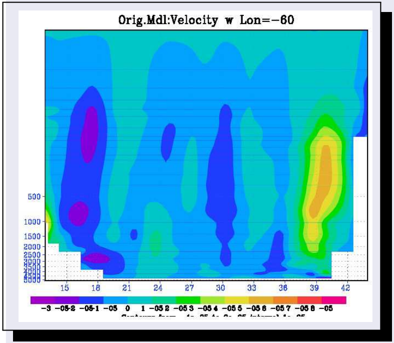
Original model



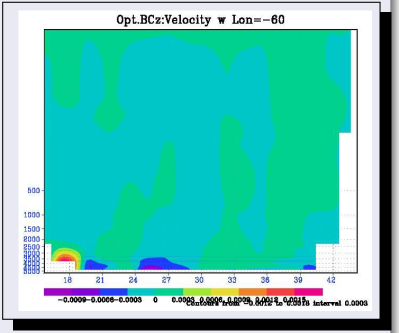
Optimal BCz



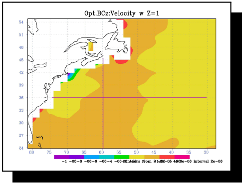
North Atlantic, January, 30, 2006, $y - z$ section



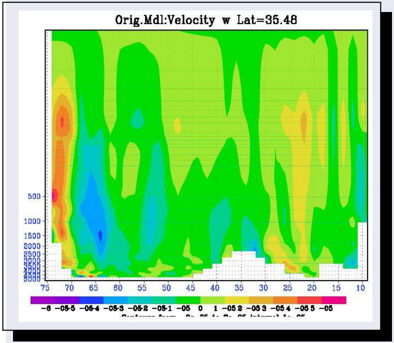
Original model



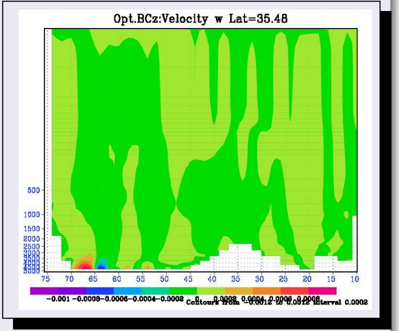
Optimal BCz



North Atlantic, January, 30, 2006, $x - z$ section

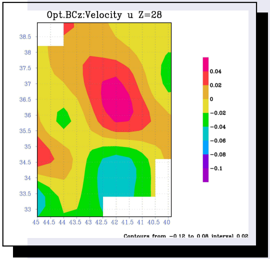


Original model

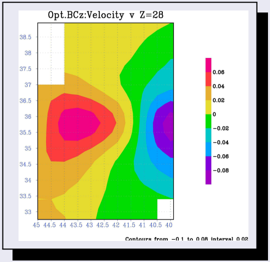


Optimal BCz

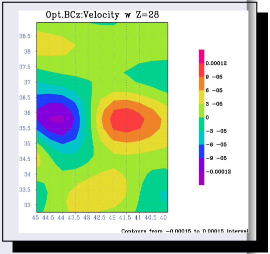
Levels $z = 28$ and $z = 29$



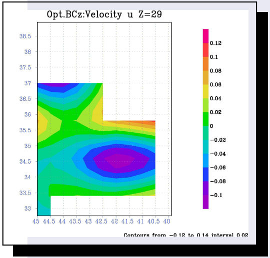
Velocity u



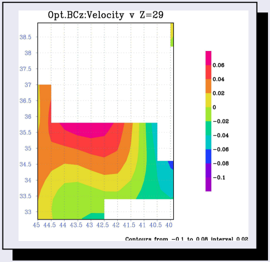
Velocity v



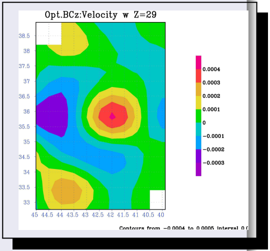
Velocity w



Velocity u



Velocity v

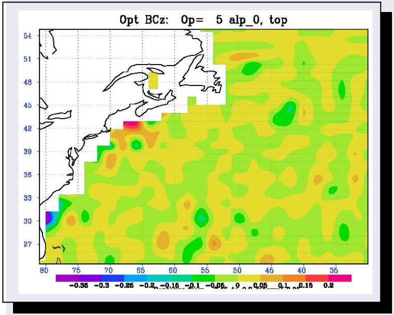


Velocity w

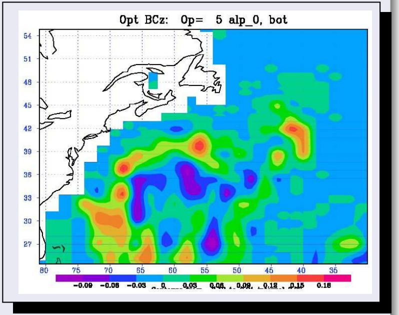
$$
\left.\frac{\partial u}{\partial z}\right|_{w_0} = \alpha_0^{DzzU^s} + \frac{\tau_x}{hz_1\rho_0}, \quad \left.\frac{\partial v}{\partial z}\right|_{w_0} = \alpha_0^{DzzU^s} + \frac{\tau_y}{hz_1\rho_0},
$$

$$
u|_{bottom} = v|_{bottom} = \alpha_0^{DzzU^b}
$$

North Atlantic

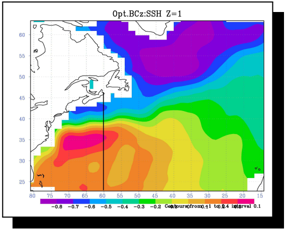


Surface

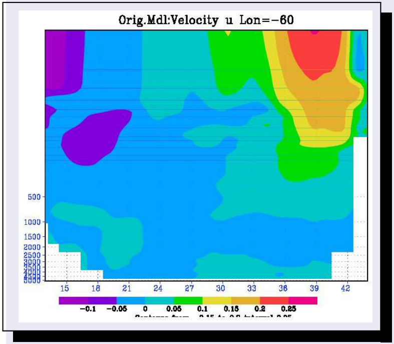


Bottom

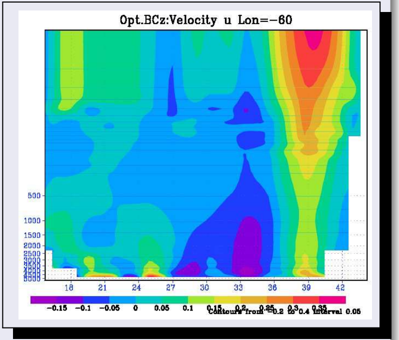
Velocity components



North Atlantic, January, 30, 2006, Velocity u, $y-z$ section



Original model



Optimal BCz

It is not an artefact.

North Atlantic

Modification of the SSH in the North Atlantic is strongly related to the boundary conditions of u and v especially on the bottom.

Only α_0 on the Bottom for u and v, only in the Vertical diffusion

$$\frac{\partial}{\partial z} A_u^z \frac{\partial u}{\partial z} \quad \text{is replaced by}$$

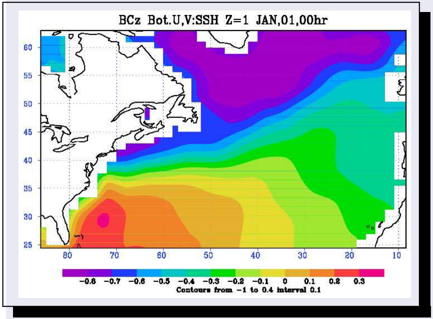
$$\begin{aligned}
\left(D_{zz}u\right)_{i,j,1/2} &= \frac{(A_u^z)_1}{hz_1 hz_{1/2}}(u_{3/2}-u_{1/2}) \\
\left(D_{zz}u\right)_{i,j,k-1/2} &= \frac{1}{hz_{k-1/2}}\left(\frac{(A_u^z)_k}{hz_k}(u_{k+1/2}-u_{k-1/2}) - \frac{(A_u^z)_{k-1}}{hz_{k-1}}(u_{k-1/2}-u_{k-3/2})\right) \quad \forall k : 2 \\
\left(D_{zz}u\right)_{i,j,K-1/2} &= \frac{1}{hz_{K-1/2}}\left[\frac{(A_u^z)_{K-1}}{hz_{K-1}}u_{K-1/2} - \left(\frac{(A_u^z)_K}{hz_K} + \frac{(A_u^z)_{K-1}}{hz_{K-1}}\right)u_{K-3/2}\right] \\
& u|_{bottom} = \alpha_0^u \quad v|_{bottom} = \alpha_0^v \qquad (2)
\end{aligned}$$

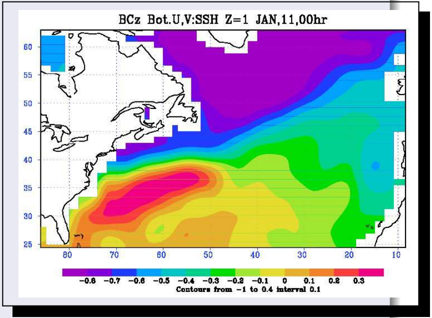
Control space dimension

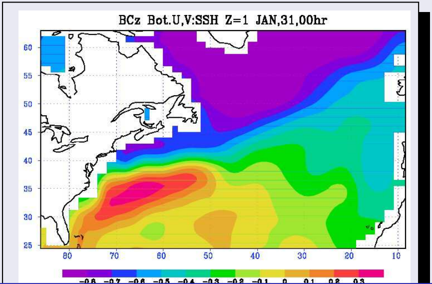
Initial conditions:	1 707 245
Full vertical boundary:	1 197 792
Only bottom:	33 272

SSH in the restrained control experiment

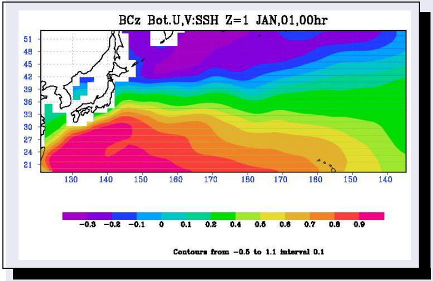
North Atlantic

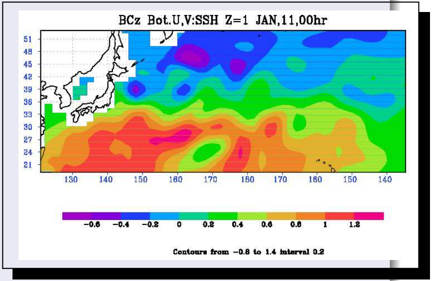


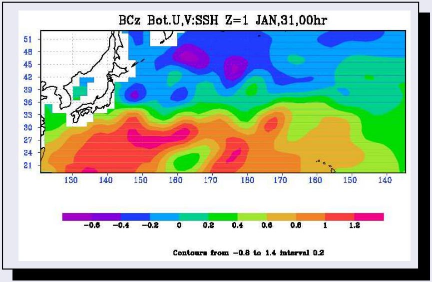




North Pacific







Extending the set of control parameters we can

- find a way to compensate model errors
- showing the most influent parameter and the most important geographical regions.

Automatic adjoint code generation helps us

- to generate TLM/AM almost immediately,
- to avoid a HUGE development/coding,
- to obtain immediately the derivative with respect to any parameter we want.

http://www-ljk.imag.fr/membres/Kazantsev/orca2/index.html